

1 Supplementary material for

2 **Supercontinent-paced magmatic destabilisation and reactivation of the Yilgarn**

3 **Craton.**

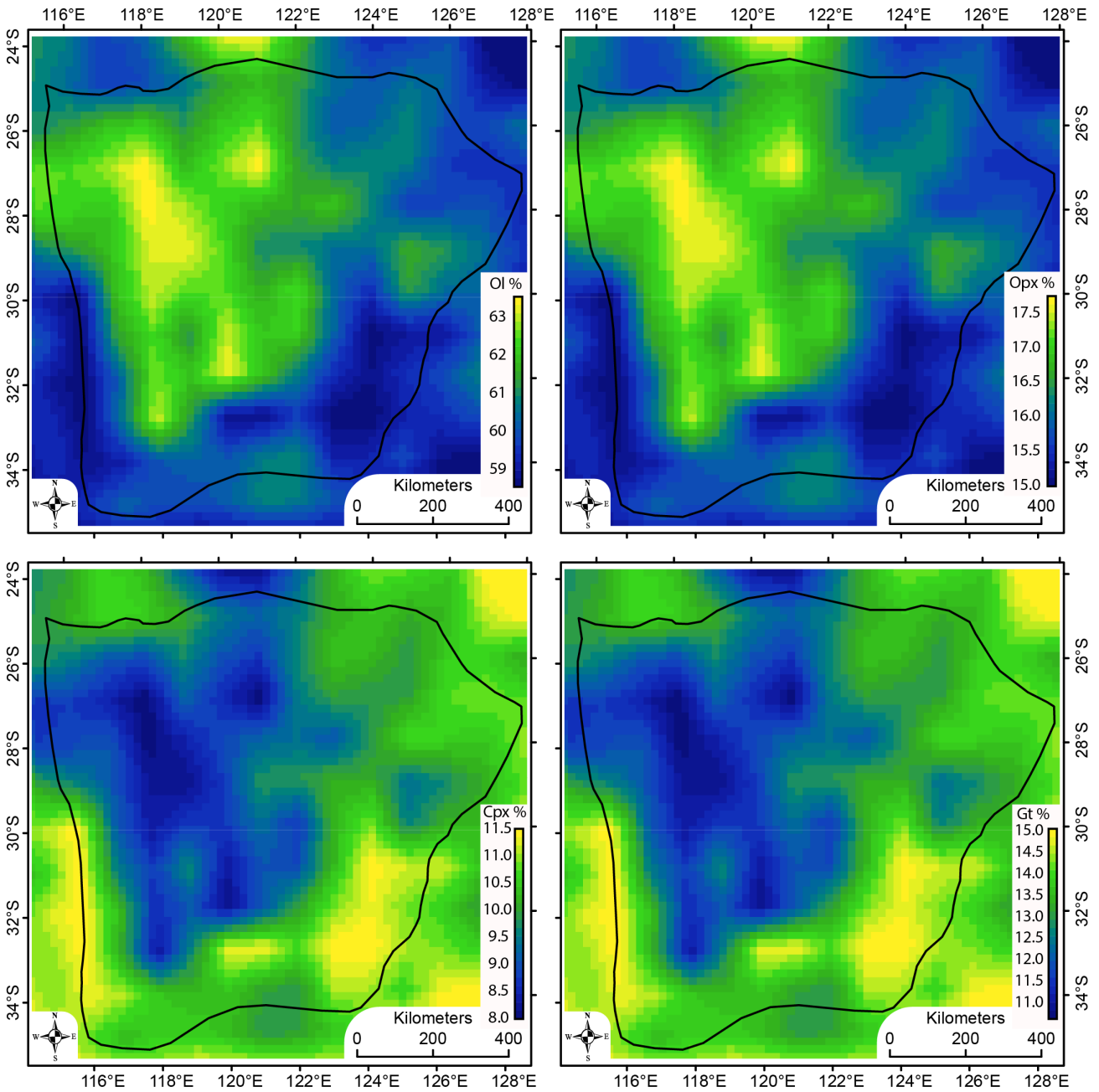
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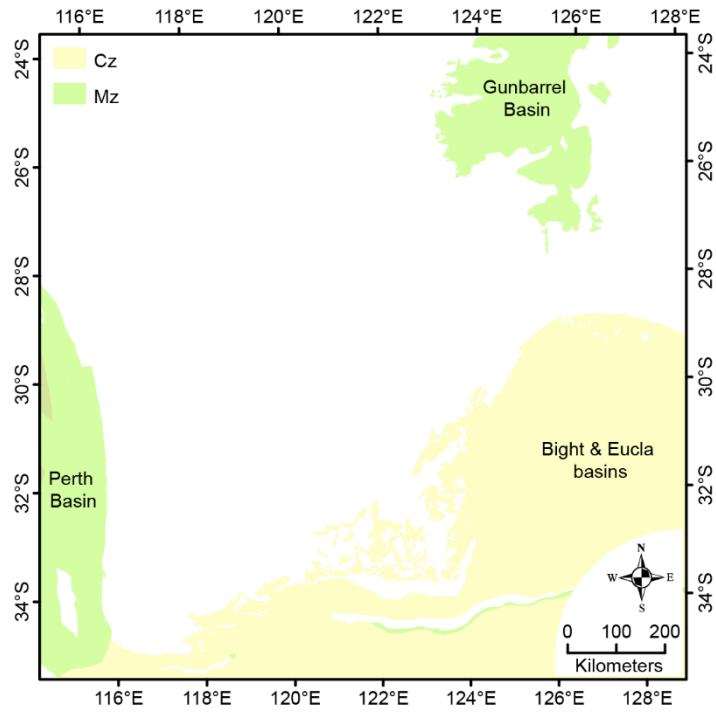
8 3- Utrecht University, Utrecht, Netherlands

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11 Figure S1: Inferred modal percentage of a) olivine, b) orthopyroxene, c) clinopyroxene, d) garnet compatible
 12 with compositional density component of the model of Tesauro et al. (2020). The model used is based on a
 13 linear interpolation between two end-member compositions, representing a 'fertile' and an Archean upper
 14 mantle, more depleted in heavy constituents (Tesauro et al., 2014).

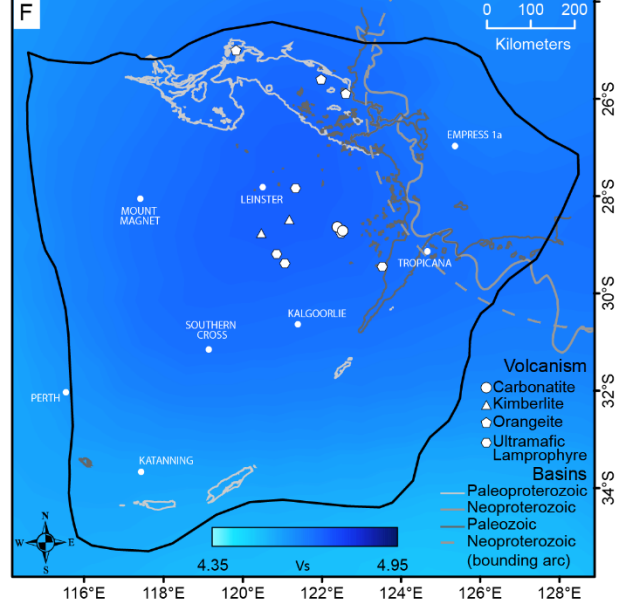
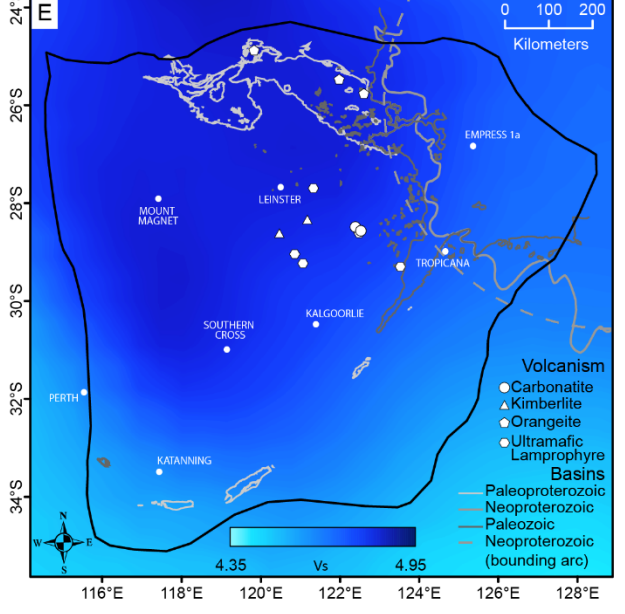
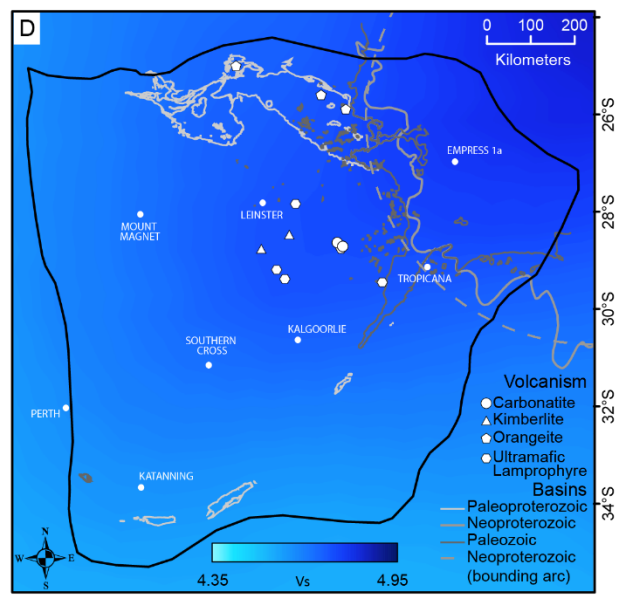
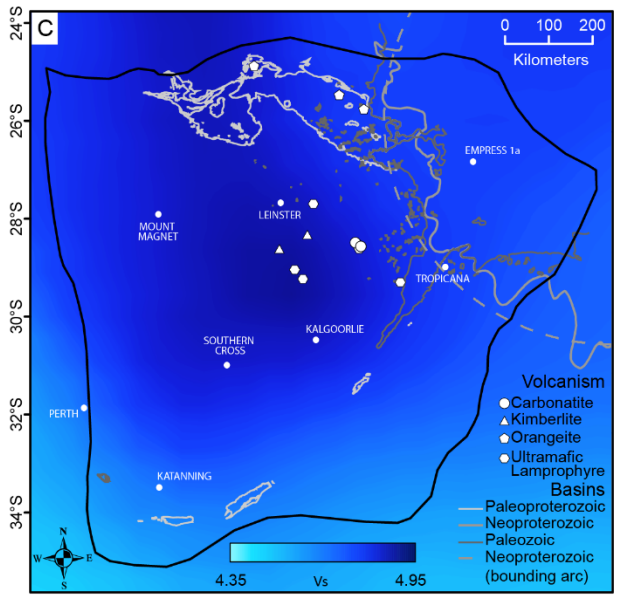
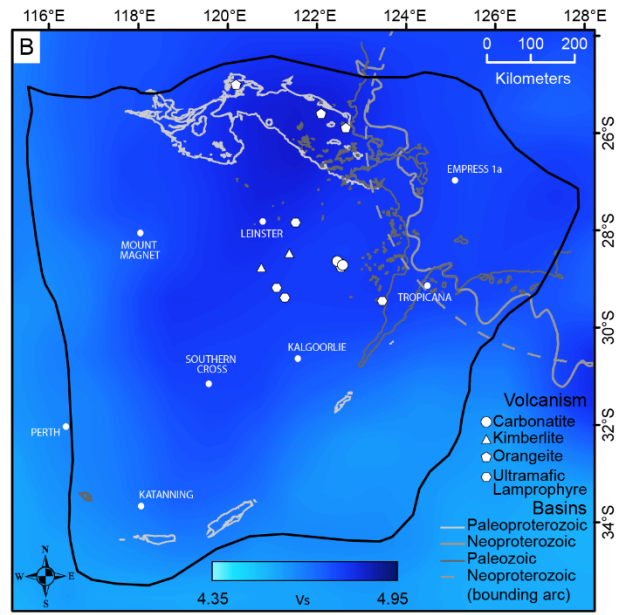
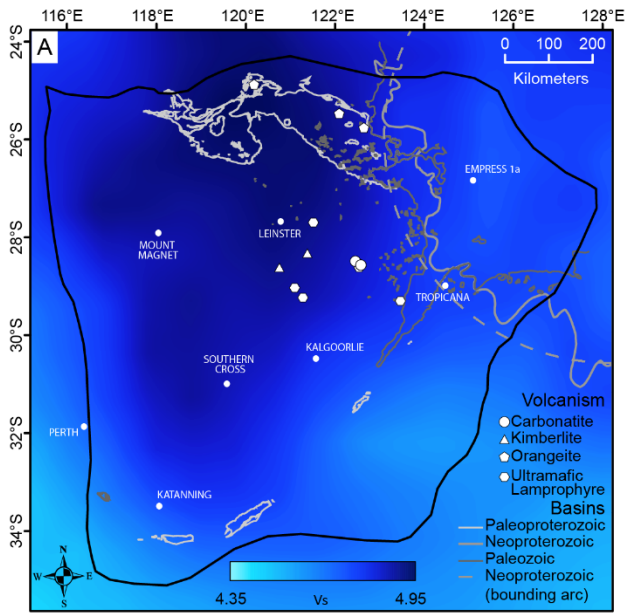


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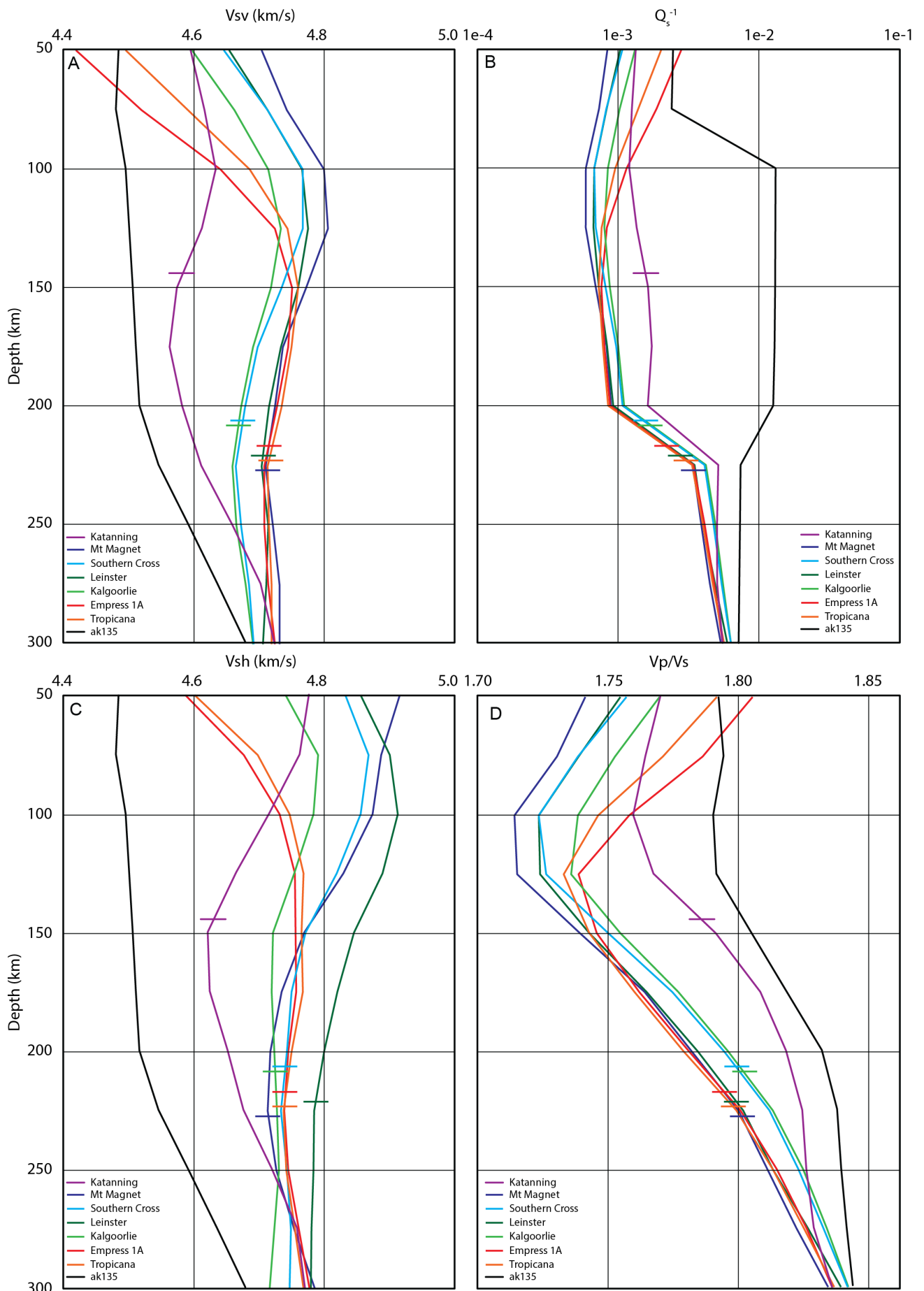
17 Figure S2: Mesozoic and Cenozoic sedimentary basins. See also slower Vs associated with Gondwana
18 breakup (Fig 4, Fig S3)

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21 Figure S3: Mantle seismic properties a) AuSREM Vsh (Kennett et al., 2013) at 100 km depth b) the same
22 model at 200 km depth c) Vsh at 100 km depth from Yoshizawa (2014) d) the same model at 200 km depth
23 e) Vsv at 100 km depth from Yoshizawa (2014) f) the same model at 200 km depth. Selected volcanic suites
24 and basin outlines are shown

25



27 Figure S4: Seismological information at seven selected sites from AuSREM (Kennett et al., 2013) and for the
28 global reference model ak135, showing a) V_{sv} , b) frequency independent Q_s^{-1} – note log scale. AuSREM
29 derives Q_s^{-1} scaled from V_s anomaly, including a scale change at 200 km depth c) V_{sh} , d) V_p/V_{sv} . Horizontal
30 bars indicate local LAB depth as the median of the models of Kennett et al. (2013), Hoggard et al. (2020)
31 after Kennett et al. (2013), Hoggard et al. (2020) after Fishwick and Rawlinson (2012), Hoggard et al. (2020)
32 after Yoshizawa (2014).